

Chemical properties of soils

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<http://timetraveler.html.xdomain.jp>

Why do we have to know the chemical properties of soils?

- **History of soil itself.**
- **Holding and supply of nutrients.**
- **Acidification, alkalization, salinization of soils.**
- **Soil pollution (organic, heavy metal pollutions)**
- **Soil improving methods.**

Understanding the soil chemical properties is necessary for the above matter.

Soil mineral components

Their change with soil
formation

Igneous rocks

Plutonic rocks



Granites



Diorite



Gabbro

Volcanic rocks



Rhyolite



Dacite



Andesite



Basalt

Rock forming minerals



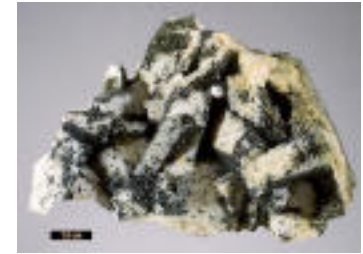
Quartz



Orthoclase



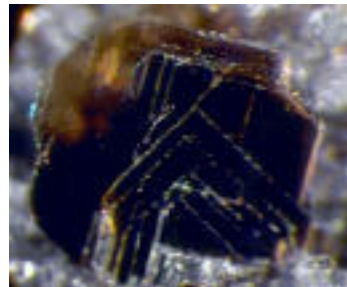
Microcline



Plagioclase



White mica



Biotite



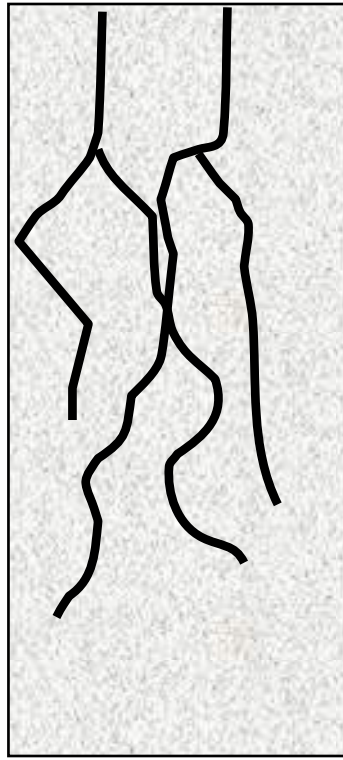
Amphibole



Pyroxene



Olivine



**Weathering
of rocks** Tens
of thousands
years

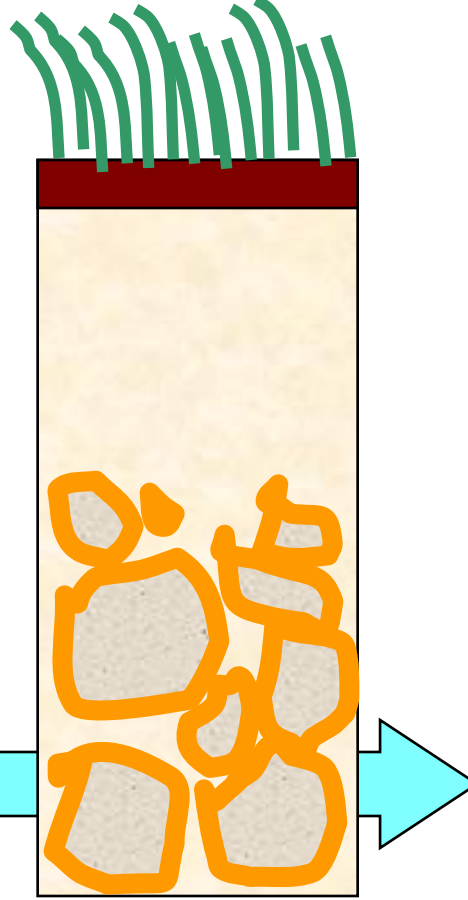
Parent
material

Climate

Biota

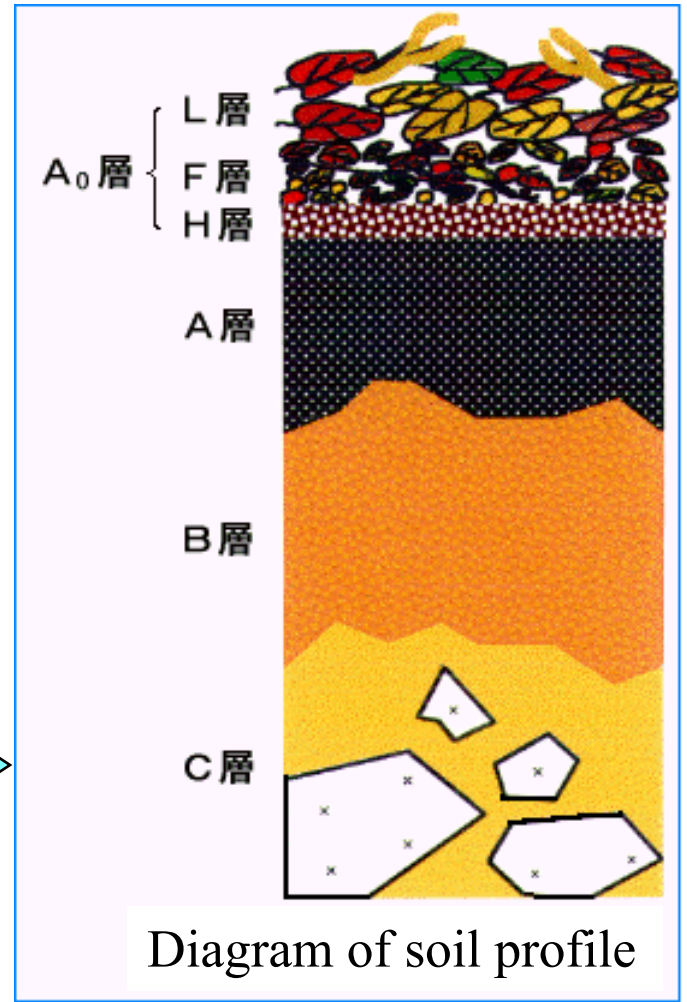
Relief

Time



Soil formation

Several
thousands years



Formation of soil layers

Formation of clay minerals

**Accumulation of soil organic
matter**

Process of soil formation

Formation of clay minerals

Formed through the reaction of rock and water near the surface of earth crust.

- **Diagenesis (deep underground, at high temperature and high pressure)**
- **Hydrothermal activity**
(Reaction with water at high temperature)
- **Weathering (at normal temperature and pressure)**

Surface area of soil particles and clay

	Maximum radius	Surface area (when packed in 1 m ³)
Gravel	10 mm	157 m ²
Coarse sand	1 mm	1570 m ²
Fine sand	0.1 mm	1.57 ha
Silt	0.01 mm	15.7 ha
Clay	0.001 mm	1.57 km ²
Kaolinite	0.05 – 0.5 μm	75 km ²
Montmorillonite	0.1 – 0.25 μm	1051 km ²
Allophane	< 0.0025 μm	1433 km ²

Specific surface area of soil particles and clay

	Maximum radius	Specific surface area (m²/g)	CEC (cmol kg⁻¹)
Gravel	10 mm	1.15×10^{-4}	
Coarse sand	1 mm	11.5×10^{-4}	
Fine sand	0.1 mm	115×10^{-4}	
Silt	0.01 mm	0.115	
Clay	0.001 mm	1.15	
Kaolinite	0.05 – 0.5 μm	55	2 - 10
Montmorillonite	0.1 – 0.25 μm	770	60 - 100
Allophane	< 0.0025 μm	1050	30 - 135

Colloid

Particles with diameter of $10^{-7} - 10^{-9}$ m dispersed in air, liquid and solid, and their dispersed state. Molecular colloid is composed of macro molecules. Particle colloid is composed of solid or liquid fine particles. Micelle colloid is composed of many associated molecules.

0.1~0.001 μm , 100~1 nm, 1000~10 \AA

Soil colloids

- **Clay minerals**
- **Humic substances**

Both are indigenous substances on the surface of earth.

Adsorb cations, anions and organic matter.

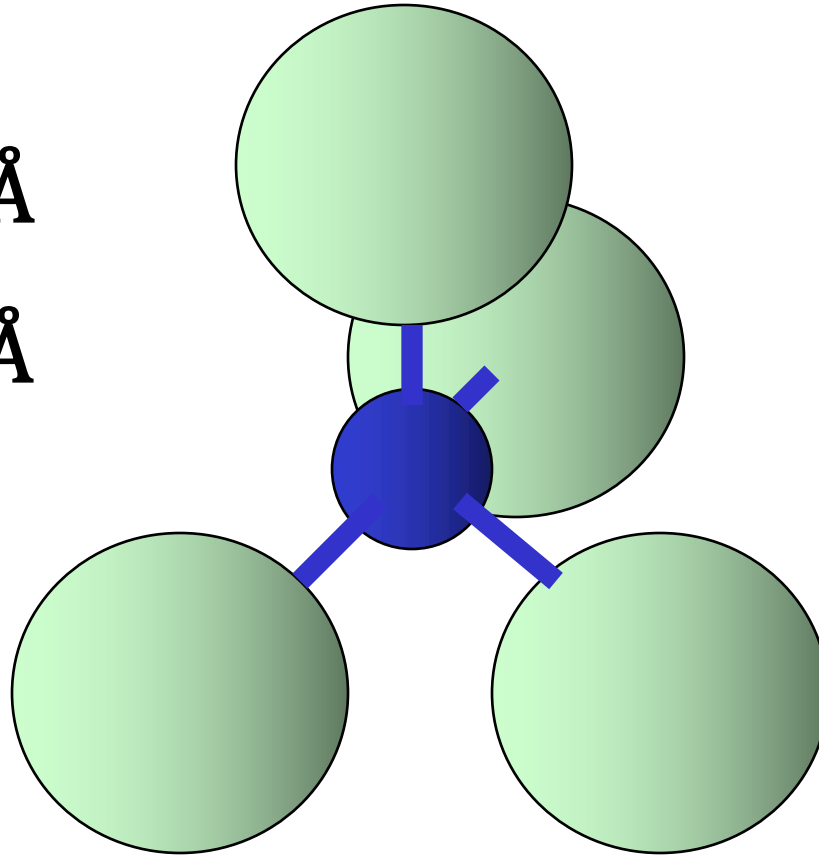
Indispensable for the soil environment suitable for life activity.

Si-O 1.62Å

Ion radius

Si^{4+} 0.40Å

O^{2-} 1.24Å



Si tetrahedron

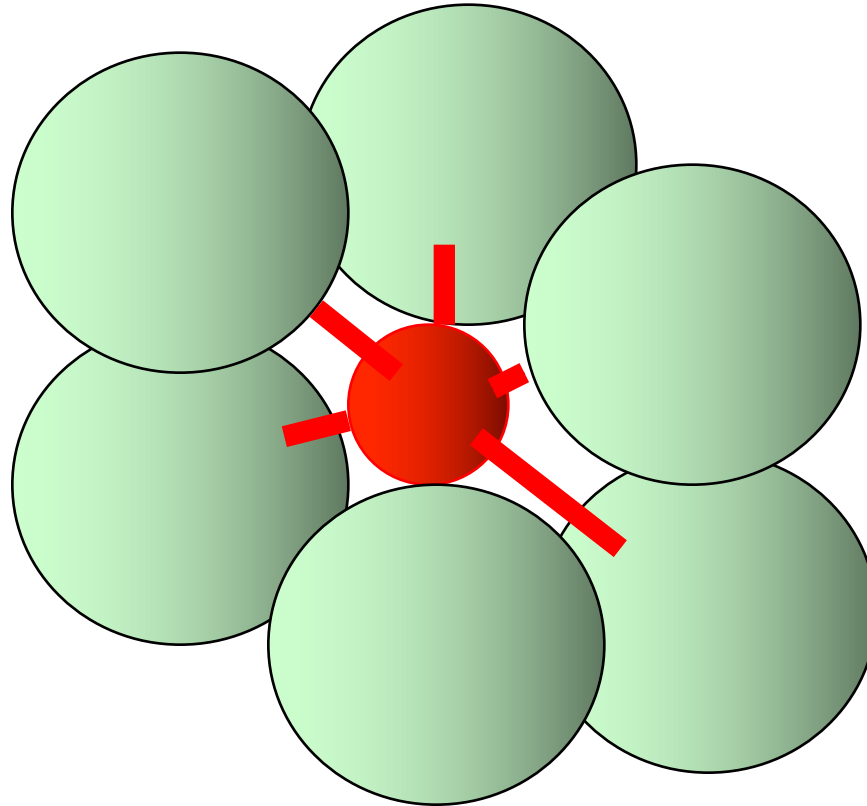
SiO₂: 50 – 60 % in rocks

Al-O 1.75 Å

Ion radius

Al^{3+} 0.51Å

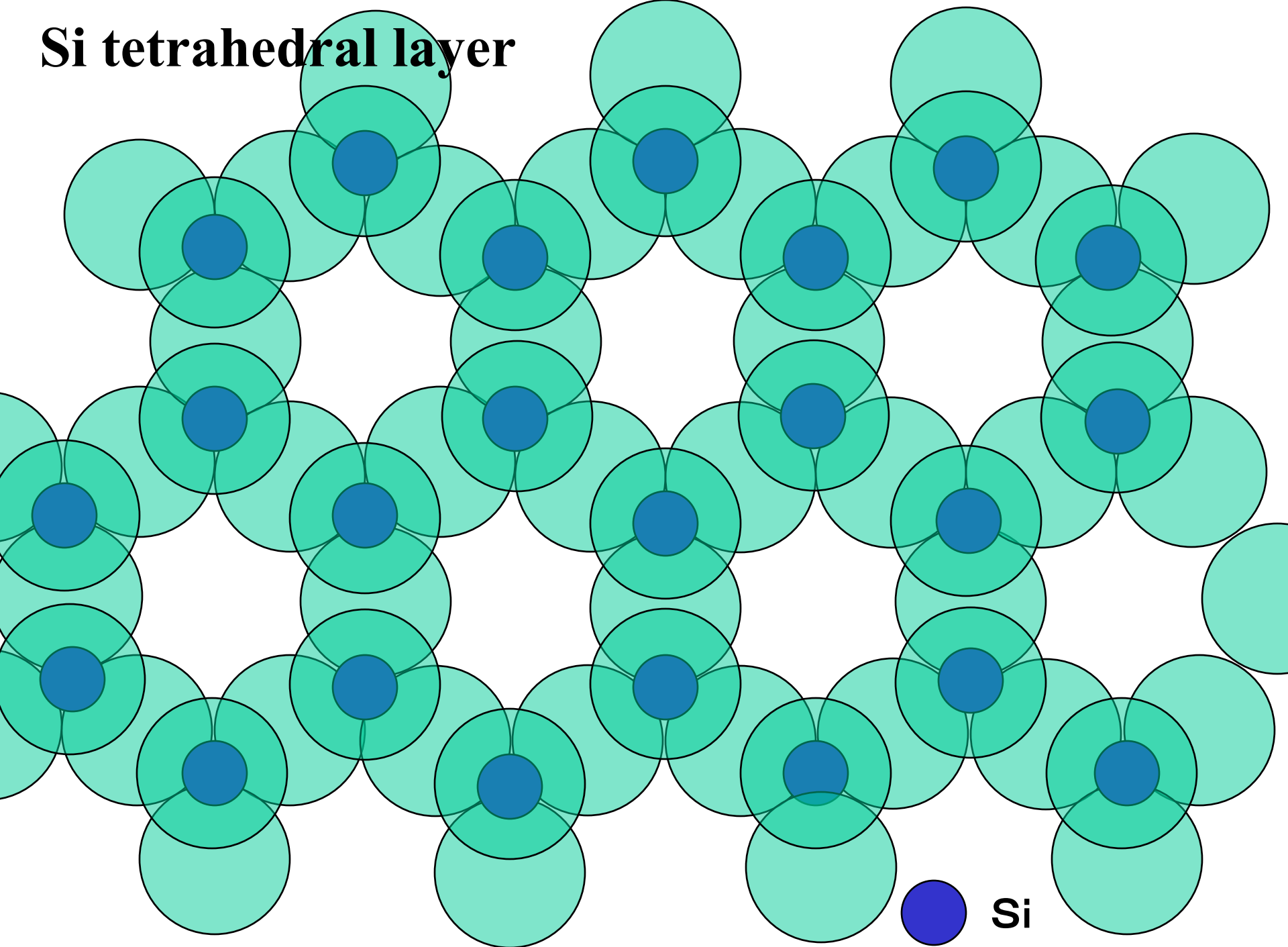
O^{2-} 1.26Å

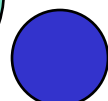


Al octahedron

Al_2O_3 : 15 – 20 % in rocks

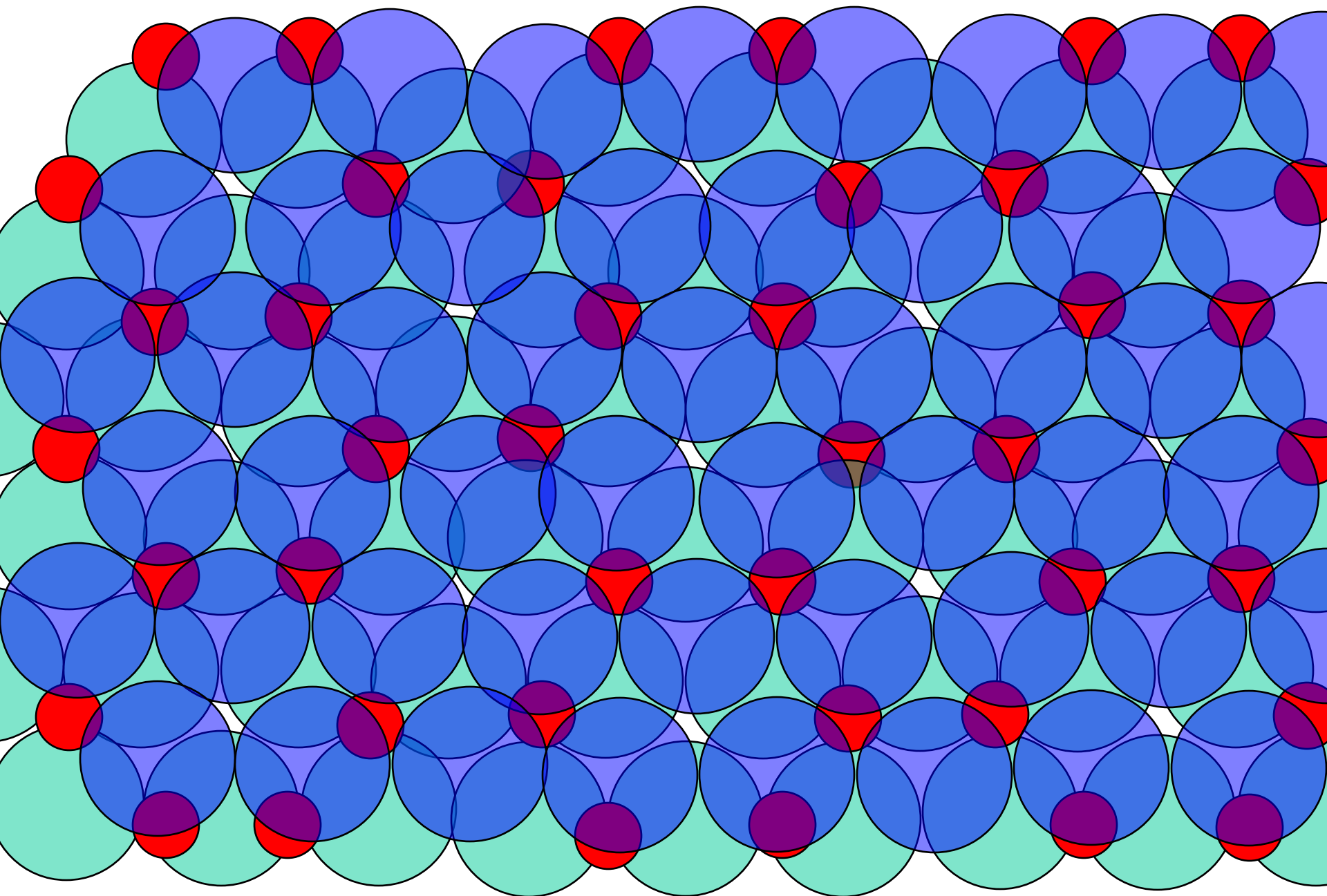
Si tetrahedral layer



 Si

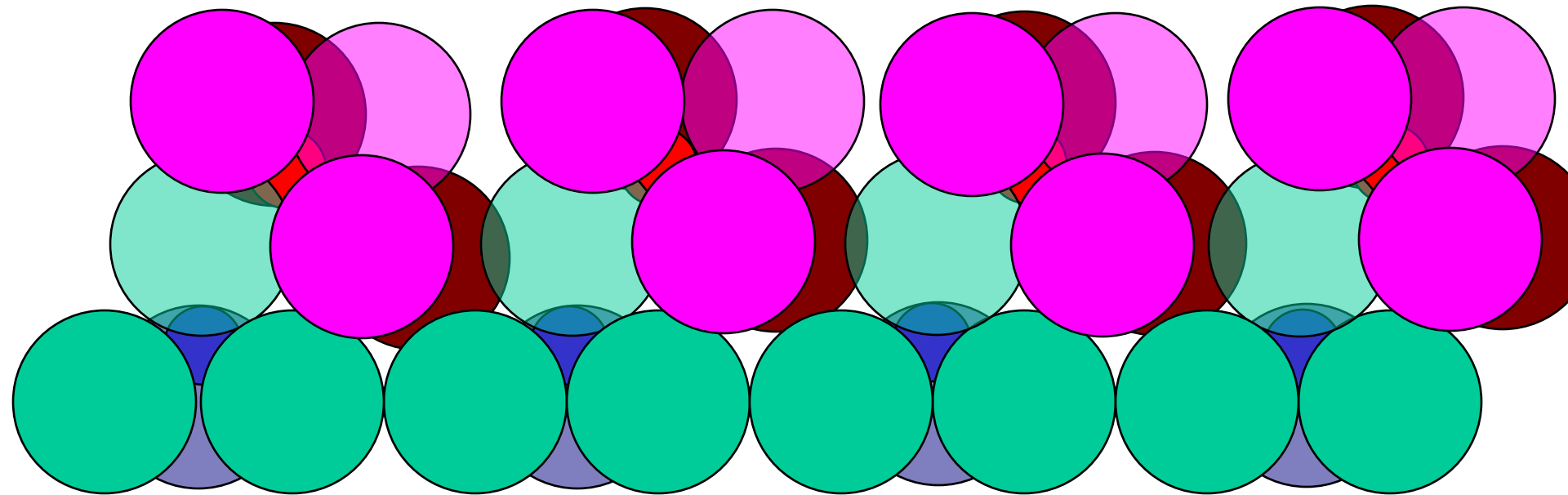
Al octahedral layer

● Al



Cross section of Al:Si 1 : 1 lattice

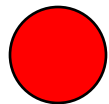
Al octahedral layer



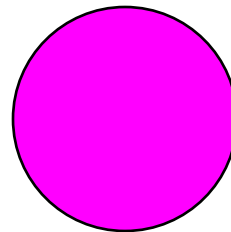
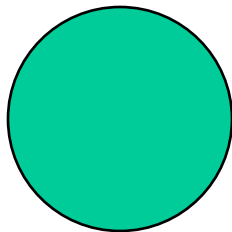
Si tetrahedral layer



Si

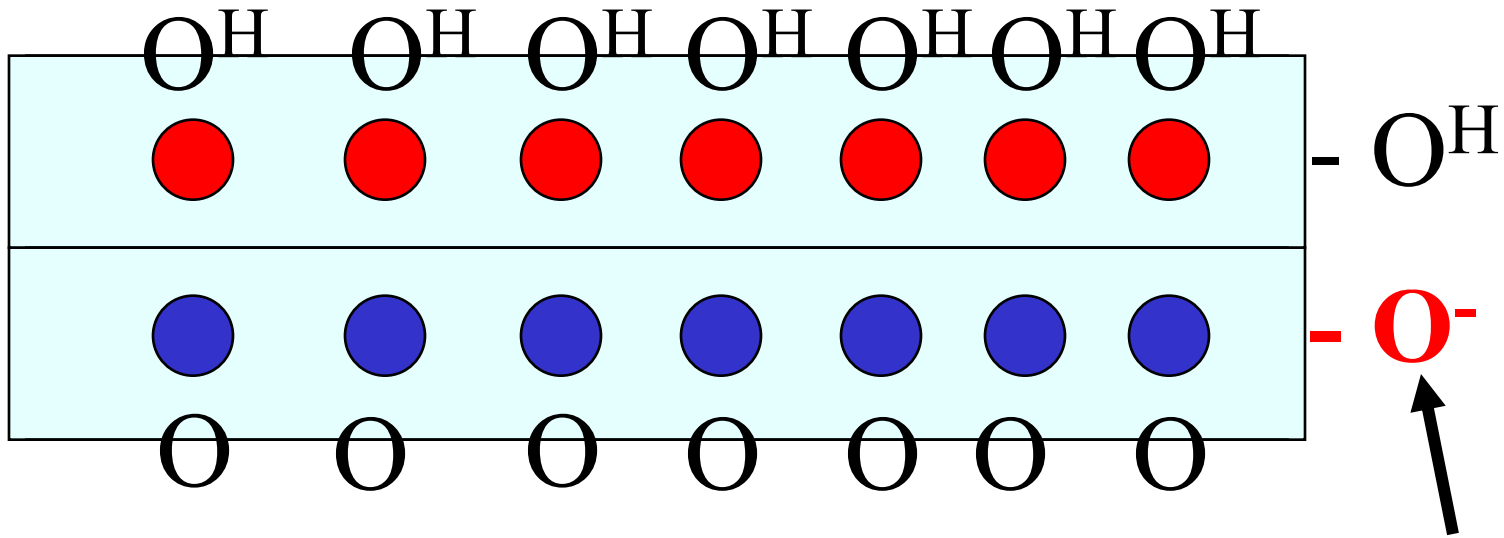


Al



Oxygen atoms

Structure of 1:1 type clay minerals



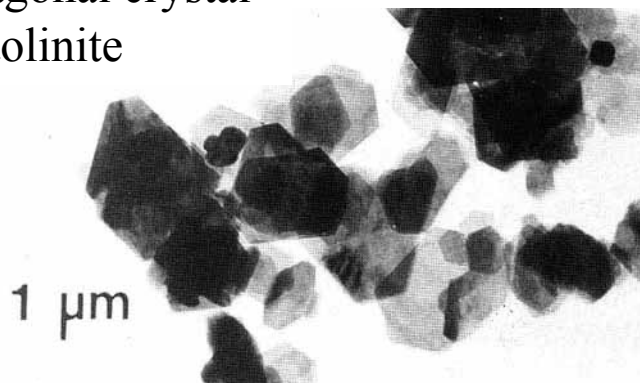
**Negative charge
on the cleaved
surface of
crystals.**

Cause of negative charge in kaolinite

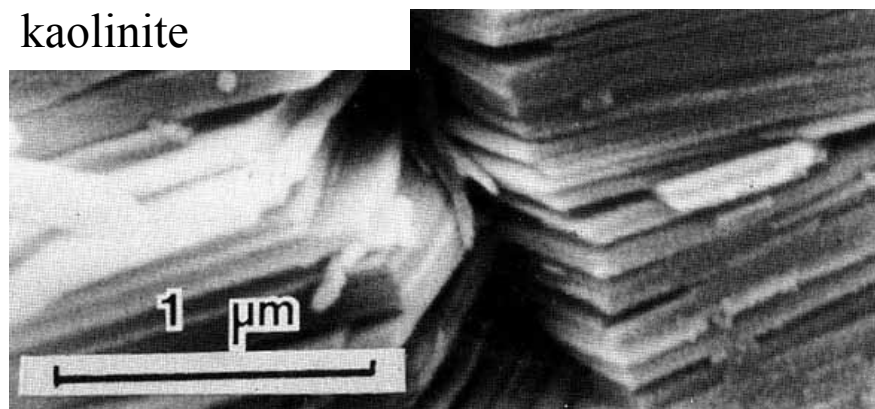
**SiO⁻ on the cleaved
surface of crystals.**



Hexagonal crystal
of kaolinite

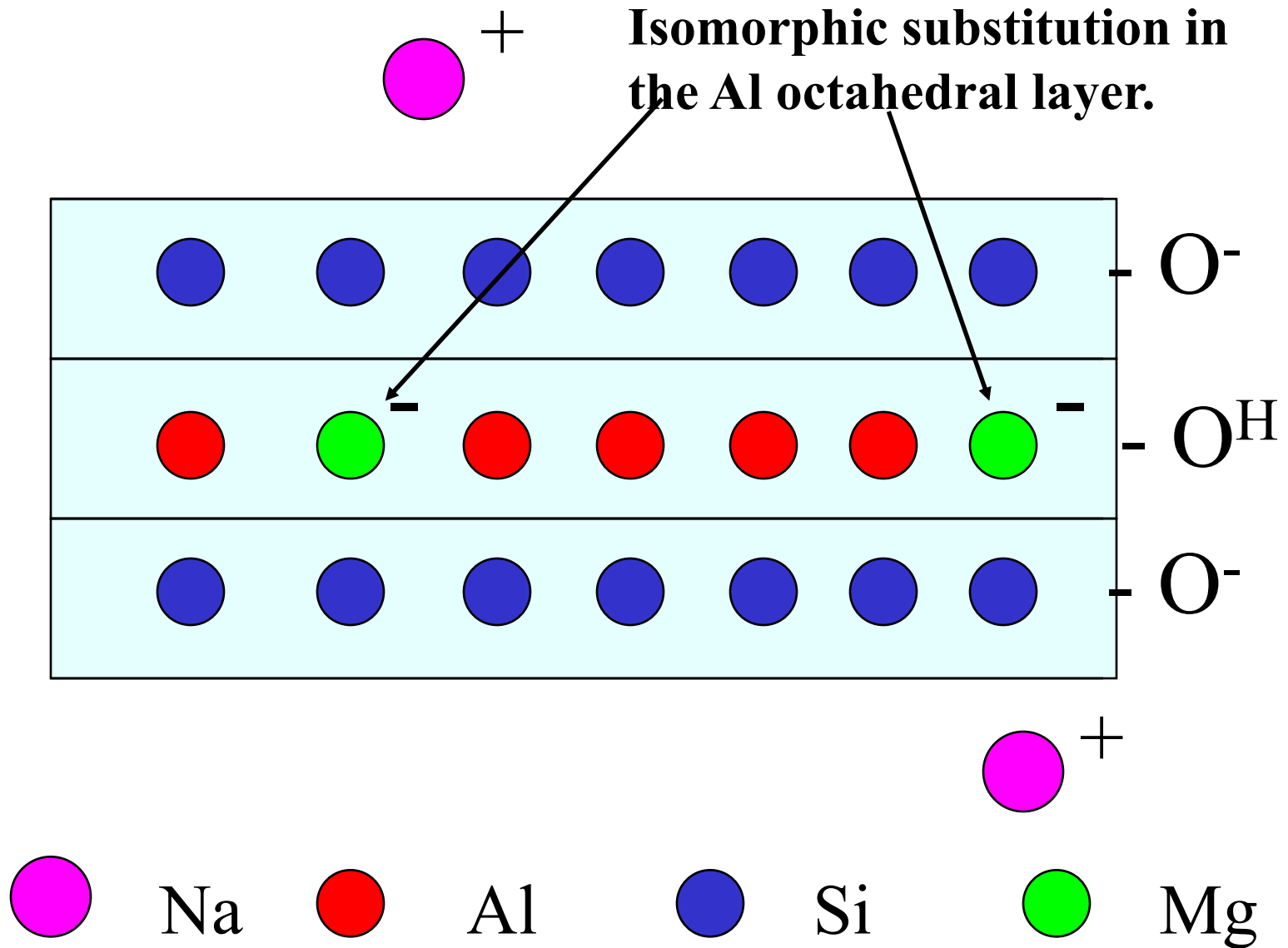


Stacked layers of
kaolinite



Kaolinite

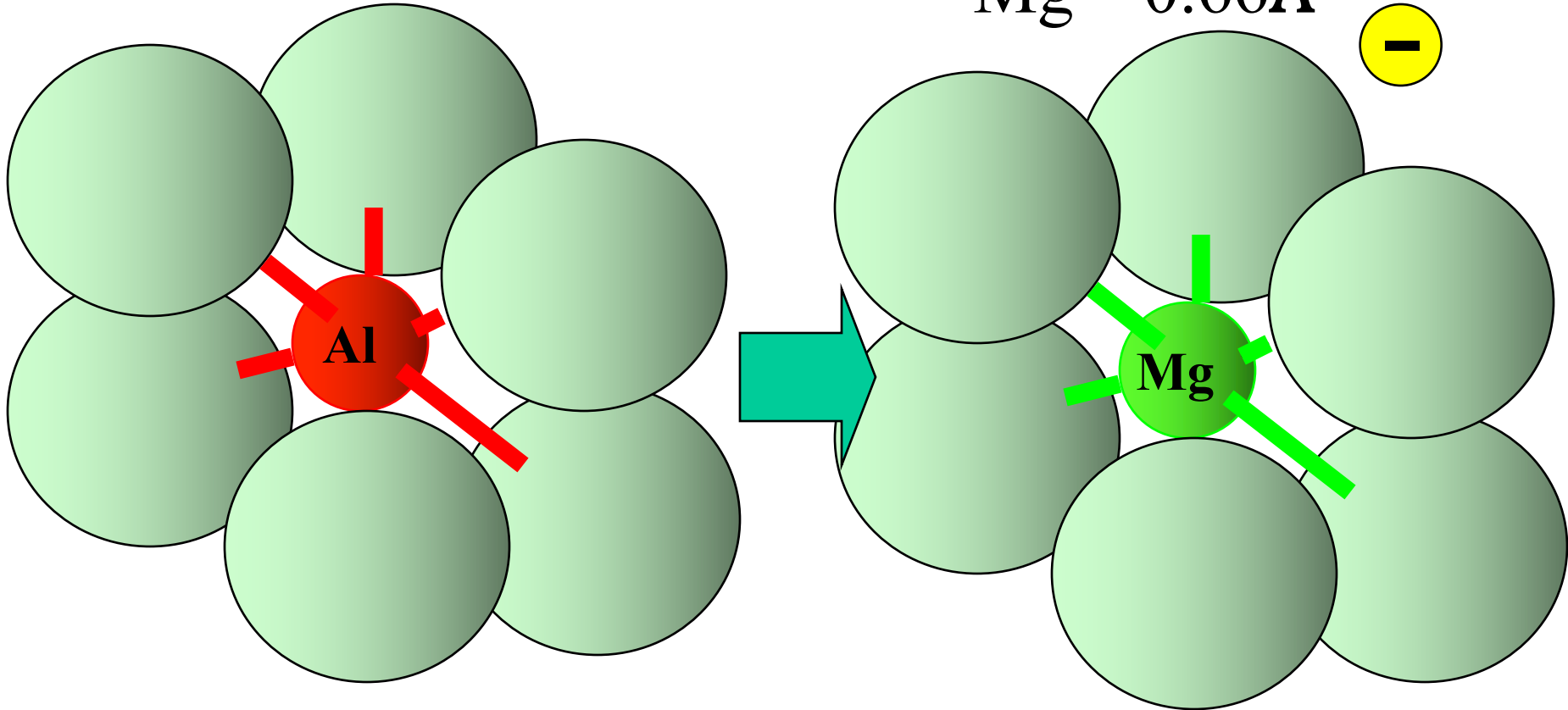
2:1 type clay minerals (eg. Montmorillonite)



Cause of negative charge in montmorillonite

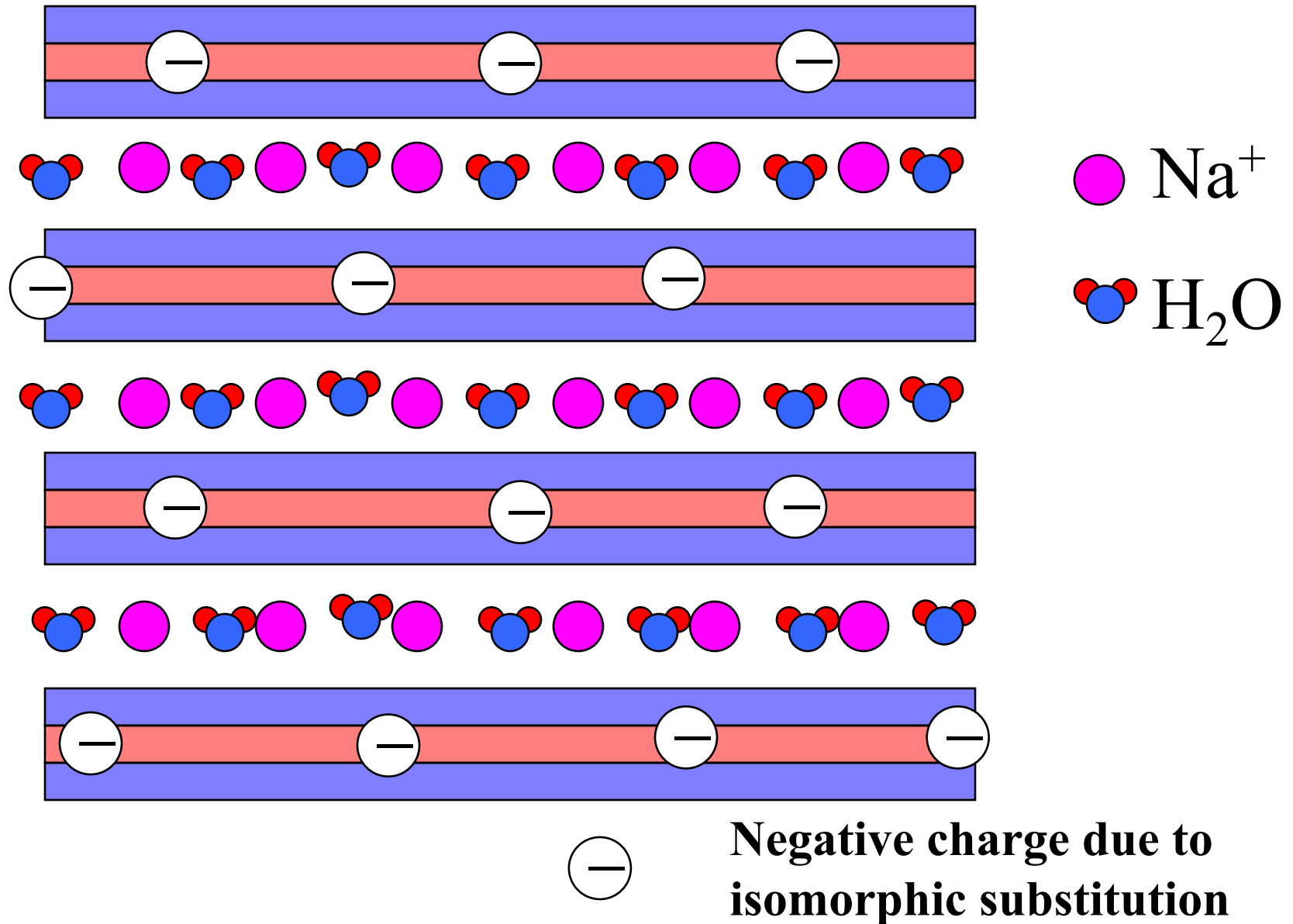
Al^{3+} 0.51Å

Mg^{2+} 0.66Å



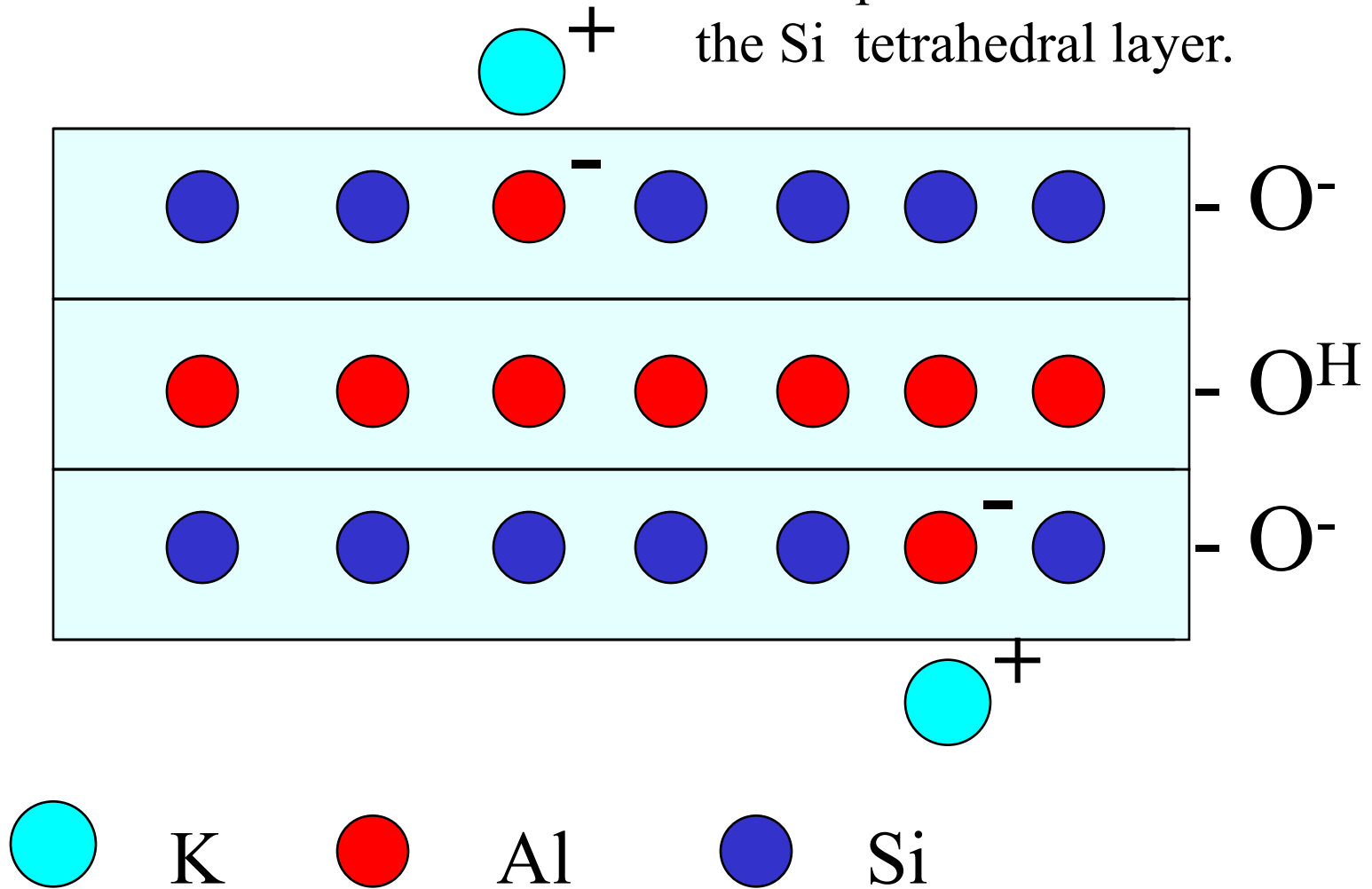
**Isomorphous substitution in the
Al octahedral layer.**

Structure of montmorillonite

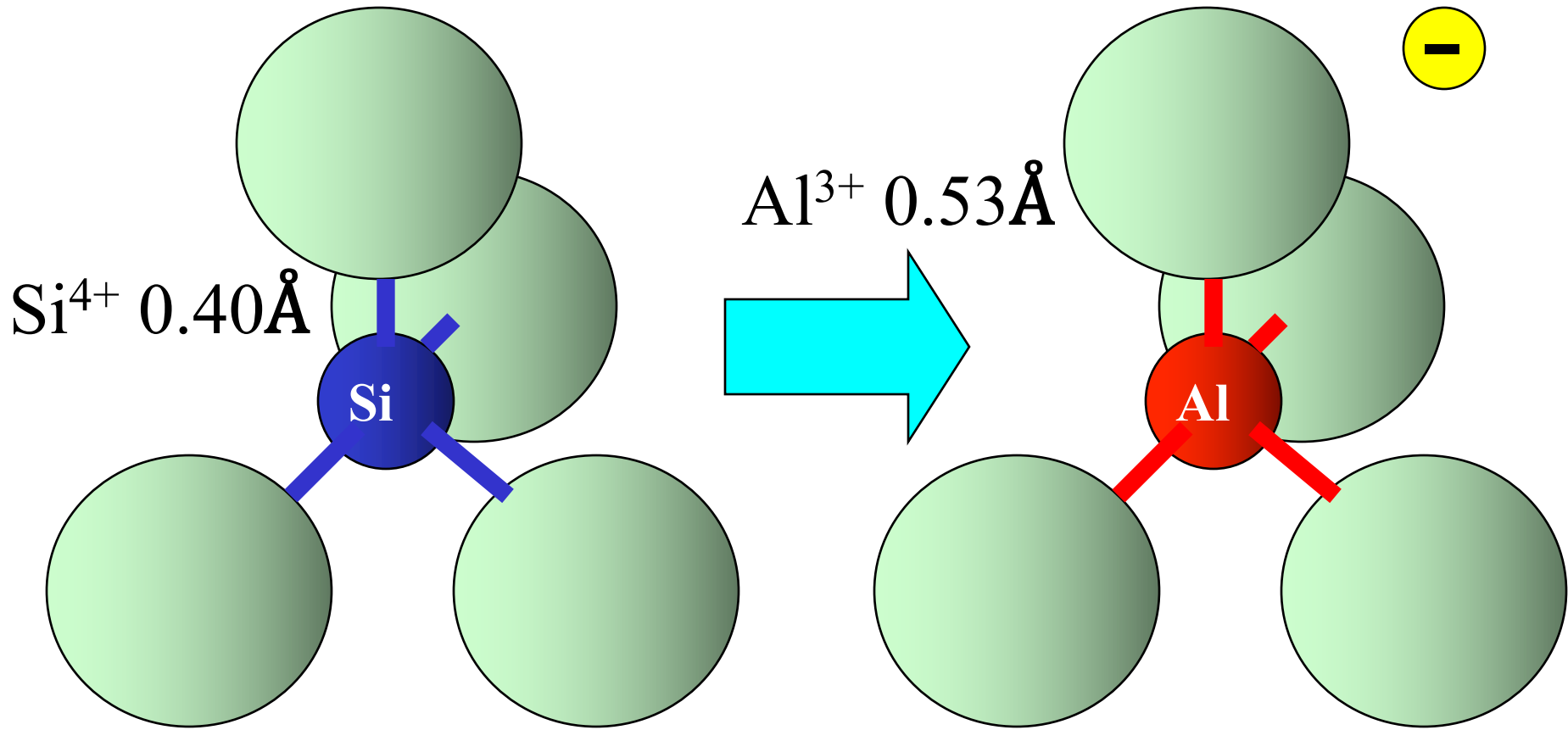


2:1 type clay minerals (Illite, Vermiculite)

Isomorphous substitution in the Si tetrahedral layer.

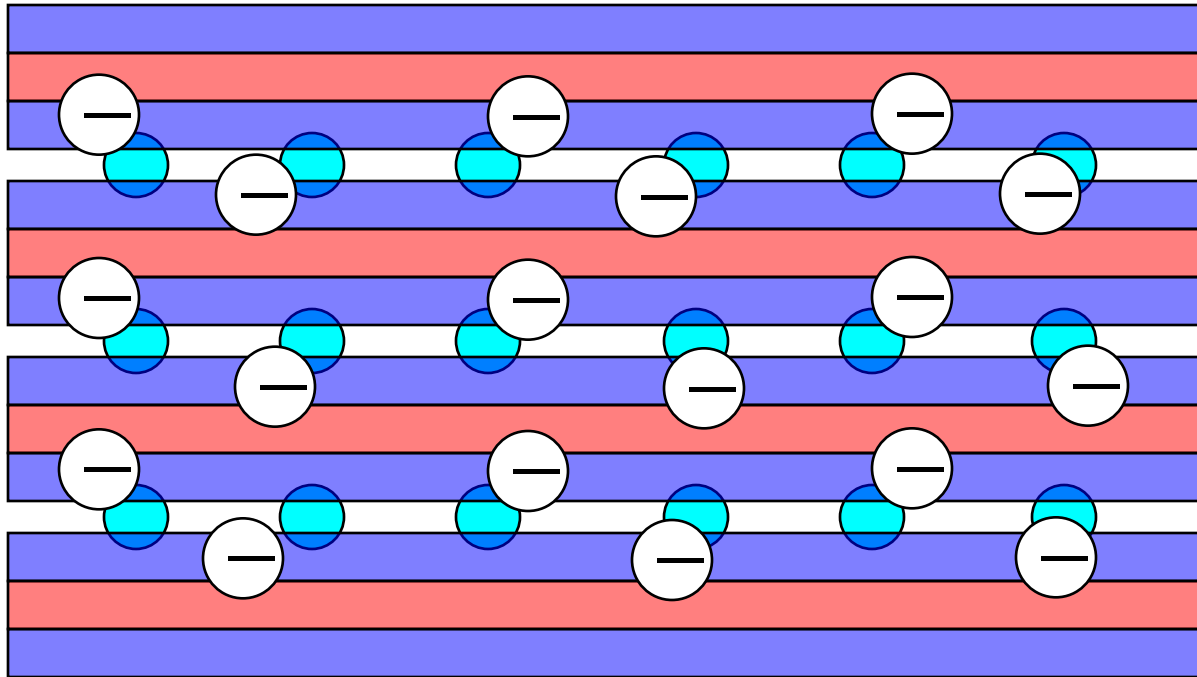


Cause of negative charge in illite.



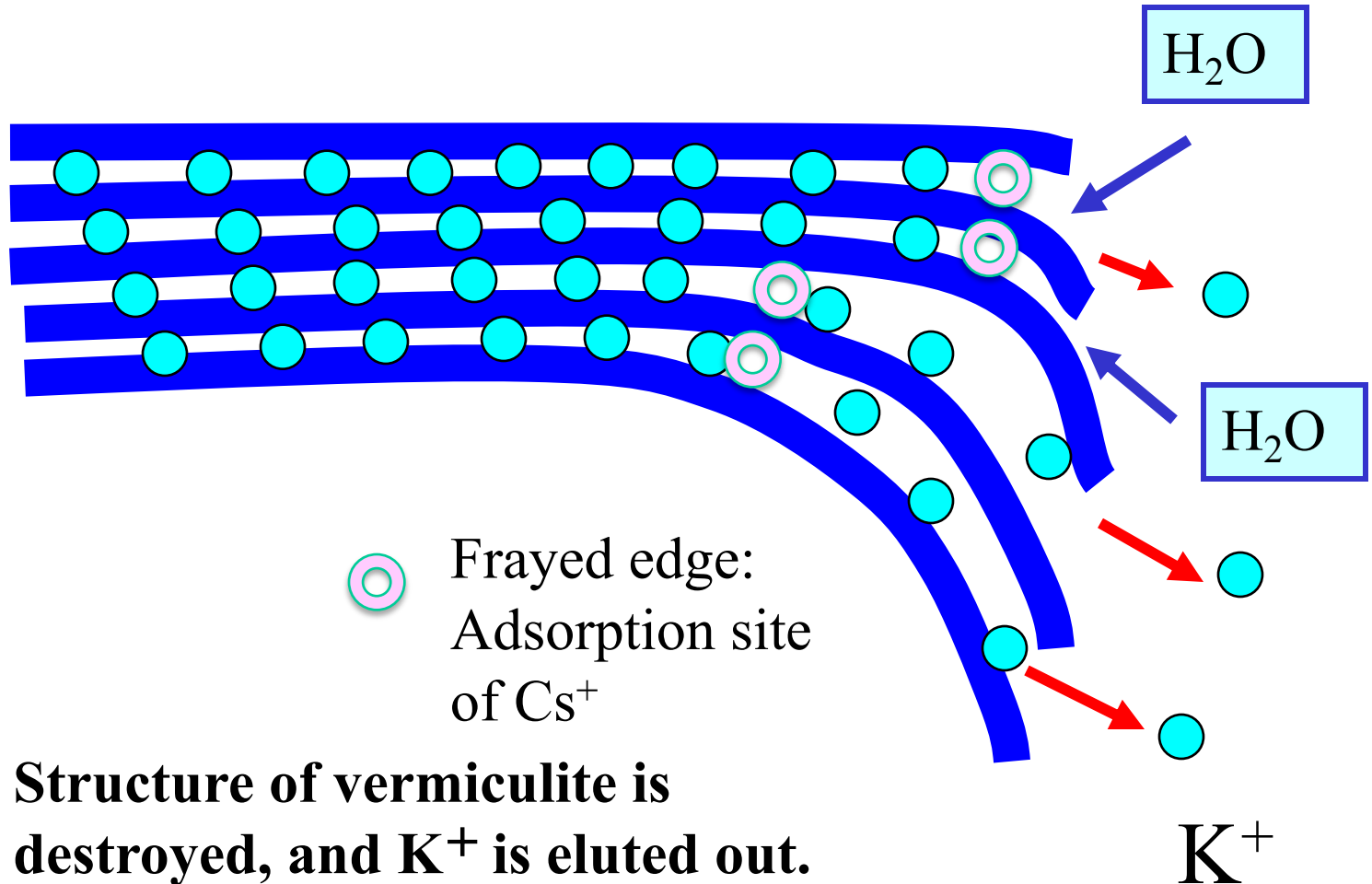
Isomorphous substitution in
the Si tetrahedral layer.

Structure of illite (mica type clay minerals)



-  K⁺
-  Negative charge due to isomorphous substitution

Structure of vermiculite



Adsorption of Cs^+ at the frayed edge

- Wedge like site in the loosened part of the stacked layers of vermiculite is called “frayed edge”.
- Hydrated cations can not enter this site.
- Unhydrated Cs^+ ion just fit this opening (niche).
- Dr. Nakao Atsushi
<http://www.kpu.ac.jp/cmsfiles/contents/0000002/2873/nakao.pdf>